

# ADAPTATION OF THE REYNOLDS STRESS TURBULENCE MODEL FOR ATMOSPHERIC SIMULATIONS

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# Outline



1. Introduction and motivations
2. Reynolds Stress model
3. Atmospheric RSM constants
4. Validation of the model
5. Conclusions and perspectives



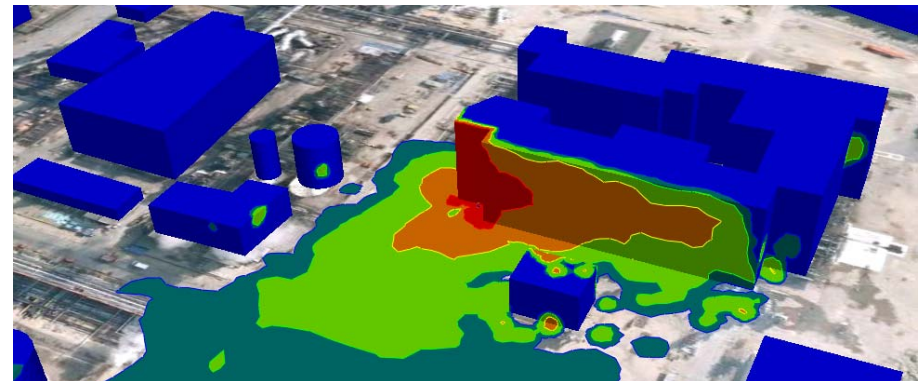
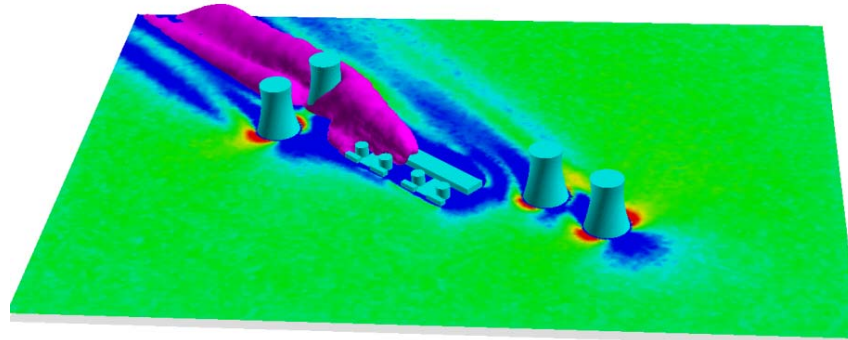


# 1 – Introduction and motivations

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## Research issues in atmospheric CFD modeling

- **Atmospheric CFD simulations (particularly RANS  $k-\epsilon$ ) are often used**
  - For local scale wind engineering
  - For dispersion in complex urban area with obstacles
  - For risk and safety assessment in industrial areas



# 1 – Introduction and motivations

## Research issues in atmospheric CFD modeling



- **Different issues have to be properly solved**
  - Boundary conditions over the surface layer?
  - Anisotropy of the turbulence?
  - “Standard” or “atmospheric/Duynkerke” constants?
  - Value of the turbulent Schmidt number?

→ Large uncertainty and user-dependent variability when comparing with field measurements
- **Objectives of this work**
  - Introduce anisotropy of turbulence using Reynolds Stress Model
  - Develop a 1D model for the entire Atmospheric Boundary Layer

→ Provide parameterizations and boundary conditions for 3D CFD calculations



## 2 – Reynolds Stress model

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### RSM equations

- **Reynolds Stress Model equations**

- Reynolds Stress equation

$$\frac{\partial \overline{u'_i u'_j}}{\partial t} + u_k \frac{\partial \overline{u'_i u'_j}}{\partial x_k} = \frac{\partial}{\partial x_k} \left( K_m \frac{\partial \overline{u'_i u'_j}}{\partial x_k} \right) + P_{ij} + \phi_{ij} - \frac{2}{3} \delta_{ij} \varepsilon$$

$$\text{with } P_{ij} = \overline{u'_i u'_k} \frac{\partial \overline{u_j}}{\partial x_k} + \overline{u'_j u'_k} \frac{\partial \overline{u_i}}{\partial x_k}$$

- Turbulent dissipation rate equation

$$\frac{\partial \varepsilon}{\partial t} + u_i \frac{\partial \varepsilon}{\partial x_i} = \frac{\partial}{\partial x_k} \left( K_m \frac{\partial \varepsilon}{\partial x_k} \right) + C_{\varepsilon 1} \frac{\varepsilon}{2k} P_{ii} - C_{\varepsilon 2} \frac{\varepsilon^2}{2k}$$

- Turbulent viscosity

$$K_m = C_\mu \frac{k^2}{\varepsilon}$$



## 2 – Reynolds Stress model

### RSM constants

- The preceding equations depends on 5 constants

$$C_{\mu}, \sigma_k, \sigma_{\varepsilon}, C_{\varepsilon 1}, C_{\varepsilon 2}$$

- Pressure-strain term  $\phi_{ij}$ 
  - We choose the model of Gibson and Launder (1978)
  - This model introduces 5 other constants :

$$C_1, C_2, C'_1, C'_2, C_L$$

- The “standard” values of these constants are:

$C_{\mu}$	$\sigma_k$	$\sigma_{\varepsilon}$	$C_{\varepsilon 1}$	$C_{\varepsilon 2}$	$C_1$	$C_2$	$C'_1$	$C'_2$	$C_L$
0.09	1.0	1.3	1.44	1.92	1.8	0.6	0.5	0.4	0.39



## 2 – Reynolds Stress model

### RSM constants



- “Standard” constants are not adapted for the atmosphere
- Consider the Surface Boundary Layer
  - For example, the  $C_\mu$  constant controls the level of turbulent kinetic energy  $k$ :

$$B = \frac{k}{u_*^2} = \frac{1}{\sqrt{C_\mu}}$$

- Wind tunnel measurements give  $B = 3.33$  (i.e.  $C_\mu = 0.09$ )
  - Atmospheric measurements (Panofsky and Dutton, 1984) give  $B = 5.48$  (i.e.  $C_\mu = 0.033$ )
- It is necessary to define “atmospheric” constants:
  - Duynkerke (1988) proposed a set of constants for the  $k$ - $\varepsilon$  model
  - In this work, we propose a new set of constants for the Reynolds Stress Model



## 3 – Atmospheric RSM constants

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## Determination approach



- **In the Atmospheric Surface Layer**
  - We assume a 1D Atmospheric Surface Layer, horizontally uniform
  - We identify the 1D RSM equations with the following analytical profiles

$$\bar{u}(z) = \frac{u_*}{\kappa} \ln\left(\frac{z}{z_0}\right)$$

$$\varepsilon(z) = \frac{u_*^3}{\kappa z}$$

$$\overline{u'^2} = \alpha_x^2 u_*^2$$

$$\overline{v'^2} = \alpha_y^2 u_*^2$$

$$\overline{w'^2} = \alpha_z^2 u_*^2$$

$$\overline{u'w'} = -u_*^2$$

$$\text{with } \begin{cases} \alpha_x = 2.46 \\ \alpha_y = 1.9 \\ \alpha_z = 1.17 \end{cases} \quad (\text{Panofsky and Dutton, 1984})$$

# 3 – Atmospheric RSM constants

## Determination approach

- It provides a system of equations for the constants:

$$\overline{u'^2} \text{ equation } -C_1 \frac{\alpha_x^2}{B} + C'_1 C_L B^{\frac{1}{2}} \kappa \alpha_z^2 - \left[ \frac{2}{3} C_2 \left( 2 - C'_2 \kappa B^{\frac{3}{2}} C_L \right) - 2 \right] \frac{\kappa z}{u_*} \frac{\partial \bar{u}}{\partial z} + \frac{2}{3} (C_1 - 1) = 0$$

$$\overline{v'^2} \text{ equation } -C_1 \frac{\alpha_y^2}{B} + C'_1 C_L B^{\frac{1}{2}} \kappa \alpha_z^2 - \frac{2}{3} C_2 \left( -1 - C'_2 \kappa B^{\frac{3}{2}} C_L \right) \frac{\kappa z}{u_*} \frac{\partial \bar{u}}{\partial z} + \frac{2}{3} (C_1 - 1) = 0$$

$$\overline{w'^2} \text{ equation } -C_1 \frac{\alpha_z^2}{B} - 2C'_1 C_L B^{\frac{1}{2}} \kappa \alpha_z^2 - \frac{2}{3} C_2 \left( -1 - 2C'_2 \kappa B^{\frac{3}{2}} C_L \right) \frac{\kappa z}{u_*} \frac{\partial \bar{u}}{\partial z} + \frac{2}{3} (C_1 - 1) = 0$$

$$\overline{u'w'} \text{ equation } \frac{1}{B} \left( C_1 + \frac{3}{2} C'_1 \right) + \left[ C_2 \left( 1 - \frac{3}{2} C'_2 \kappa B^{\frac{3}{2}} C_L \right) - 1 \right] \alpha_z^2 \frac{\kappa z}{u_*} \frac{\partial \bar{u}}{\partial z} = 0$$

- Which provides, after resolution, a set of “atmospheric” constants:

$C_\mu$	$\sigma_k$	$\sigma_\epsilon$	$C_{\epsilon 1}$	$C_{\epsilon 2}$	$C_1$	$C_2$	$C'_1$	$C'_2$	$C_L$
0.033	1.0	2.38	1.46	1.83	1.8	0.6	0.94	0.03	0.19



## 4 – Validation of the model

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### Methodology of the 1D numerical model



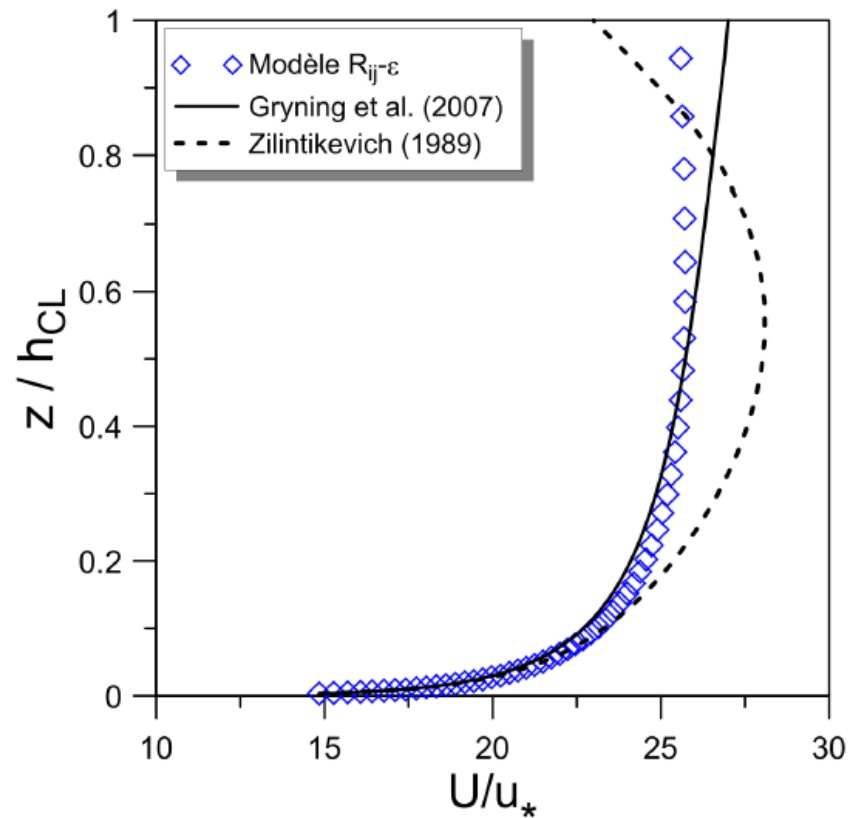
- **We have developed a numerical 1D model for the ABL**
  - Flow is horizontally homogenous:
$$\bar{w} = 0 \quad \text{and} \quad \frac{\partial}{\partial x} = \frac{\partial}{\partial y} = 0$$
  - Pressure gradient and Coriolis force in geostrophic balance
  - 1D equation model of mean wind speed, including Coriolis effect:
$$\frac{\partial \bar{u}}{\partial t} = -\frac{\partial \overline{u'w'}}{\partial z} + f(\bar{v} - v_g) \quad \text{and} \quad \frac{\partial \bar{v}}{\partial t} = -\frac{\partial \overline{v'w'}}{\partial z} - f(\bar{u} - u_g)$$
  - Reynolds Stress Model as turbulence closure model
  - Equations are solved numerically until a steady state
- **We apply the model on the overall ABL and compare it with empirical results**

## 4 – Validation of the model

### Results in neutral conditions

- **Velocity profile:**

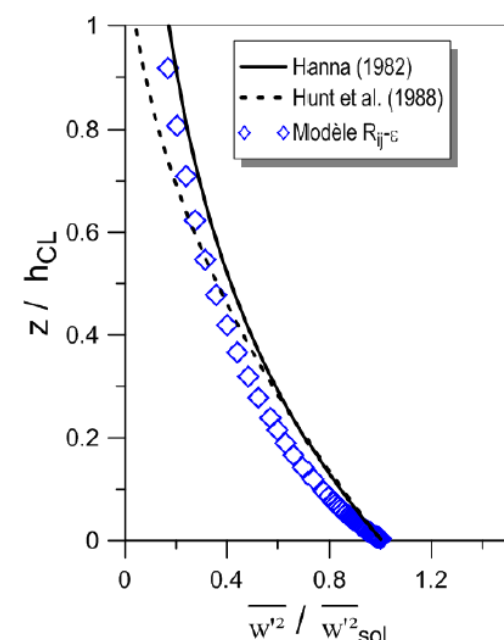
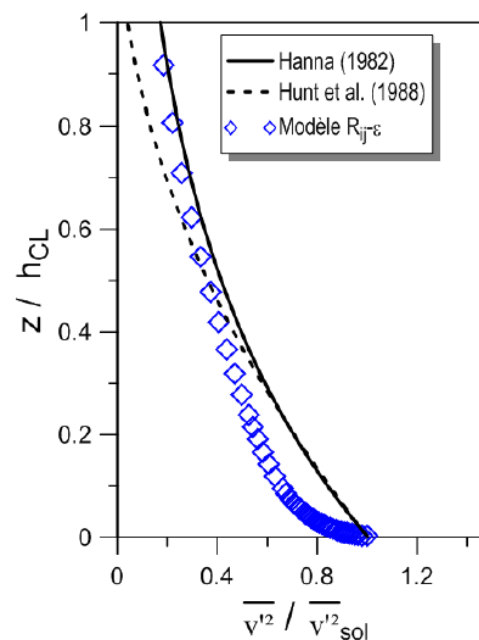
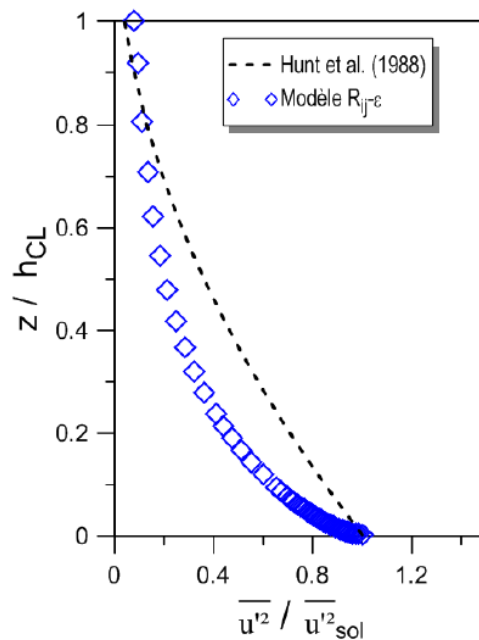
- For example for  $U_g = 5\text{m}\cdot\text{s}^{-1}$ ,  $z_0 = 0.01\text{m}$ ,  $\varphi = 45^\circ$



# 4 – Validation of the model

## Results in neutral conditions

- **Turbulence profiles (normalized by the ground value):**
  - For example for  $U_g = 5\text{m}\cdot\text{s}^{-1}$ ,  $z_0 = 0.01\text{m}$ ,  $\varphi = 45^\circ$





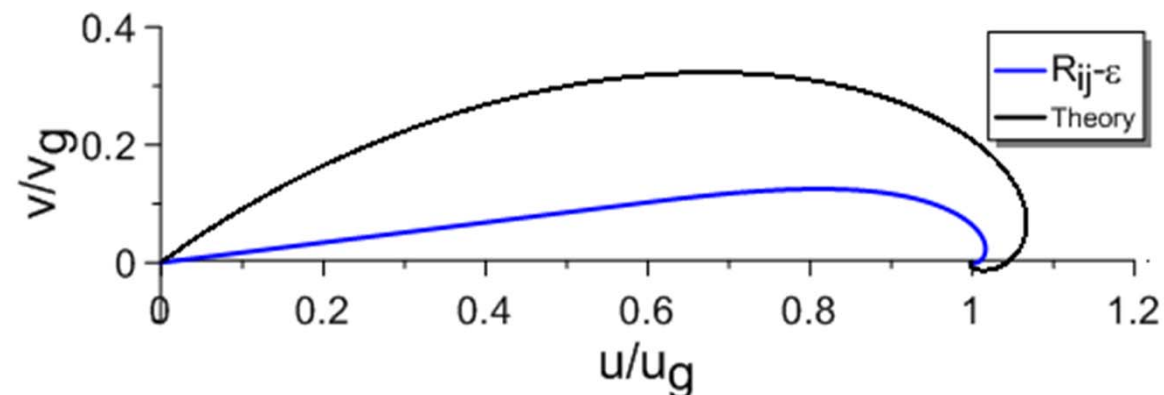
## 4 – Validation of the model

### Results in neutral conditions

- Ekman's theory predicts the twisting of the flow in the Ekman layer:

$$\begin{cases} u = U_g [1 - \exp(-az) \cos(az)] \\ v = \text{sgn}(f) U_g \exp(-az) \sin(az) \end{cases} \quad \text{with} \quad a = \sqrt{\frac{|f|}{2K_M}}$$

- Simulation of the Ekman Layer with the 1D model:



## 4 – Validation of the model

### Results in neutral conditions

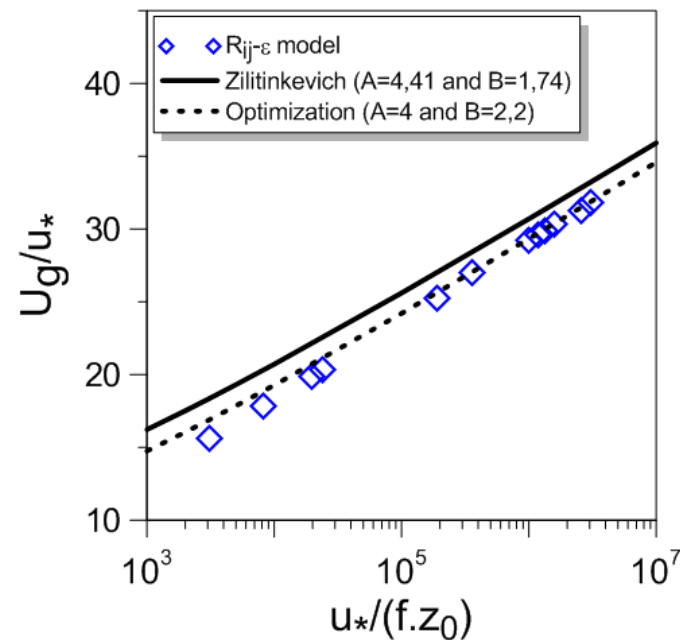
- The Rossby similarity theory gives:

$$\frac{U_g}{u_*} = \frac{1}{\kappa} \sqrt{\left( \ln \left( \frac{u_*}{f z_0} \right) - B \right)^2 - A^2}$$

*A and B are empirical constants*

- Sensibility of the 1D model to control parameters

$U_g$ ,  $f$  and  $z_0$

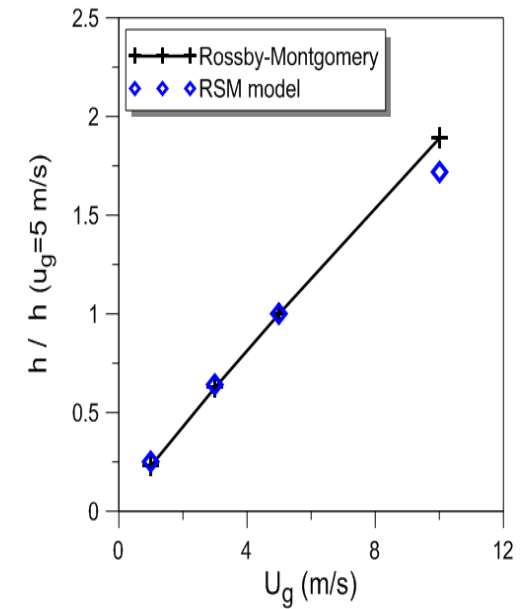
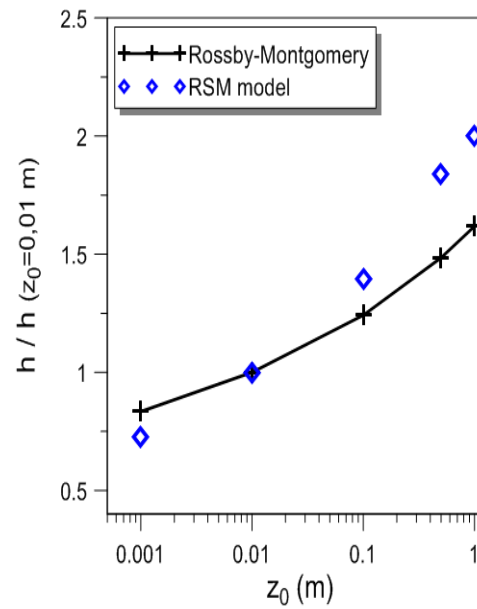
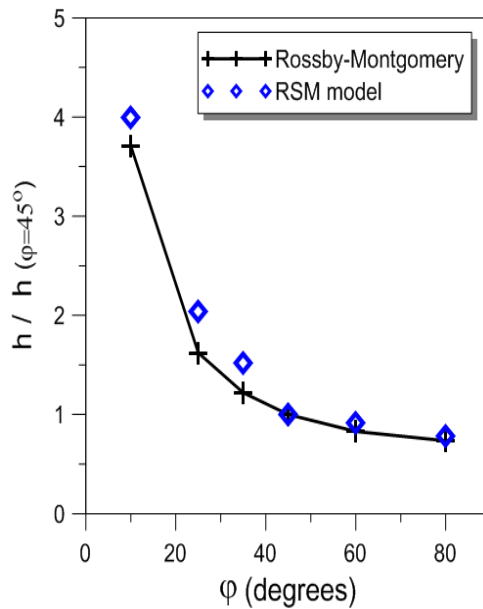


# 4 – Validation of the model

## Results in neutral conditions

- **Atmospheric Boundary Layer height:**
  - The Rossby-Montgomery equation

$$h = \frac{c u_*}{f} \quad \text{with } c \approx 0.26$$

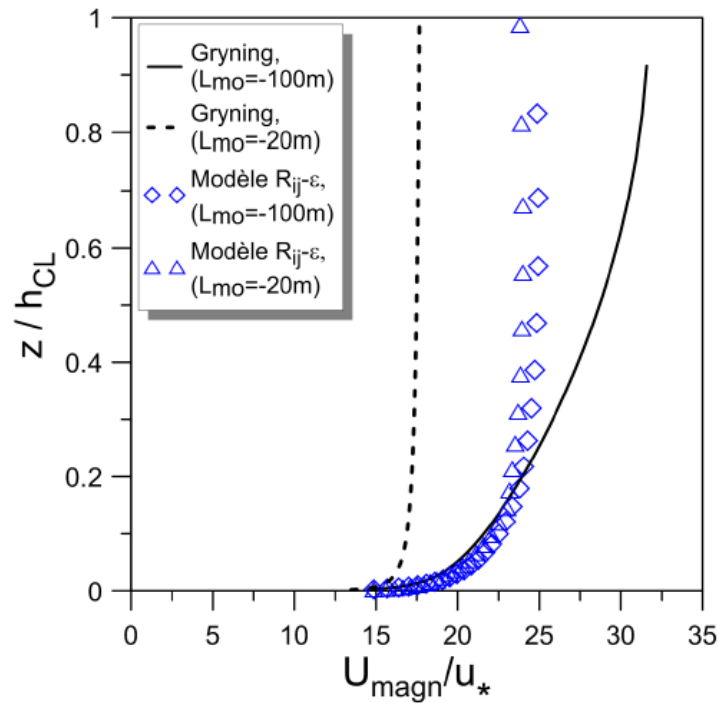


# 4 – Validation of the model

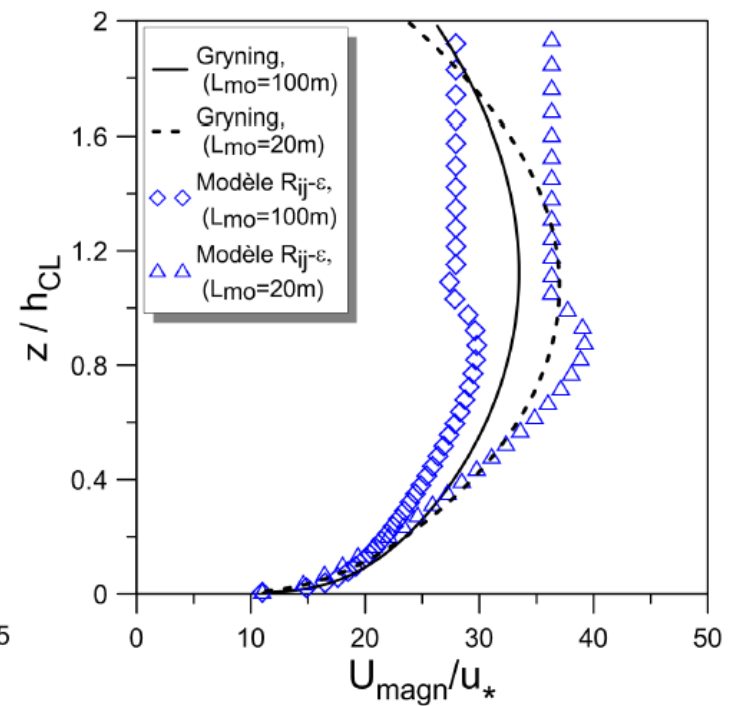
## Results in stratified conditions

- **Velocity profile:**

- For example for  $U_g = 5\text{m}\cdot\text{s}^{-1}$ ,  $z_0 = 0.01\text{m}$ ,  $\varphi = 45^\circ$



Unstable conditions

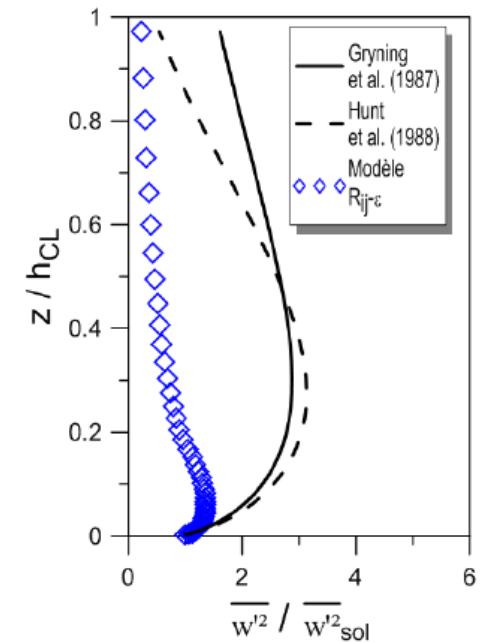
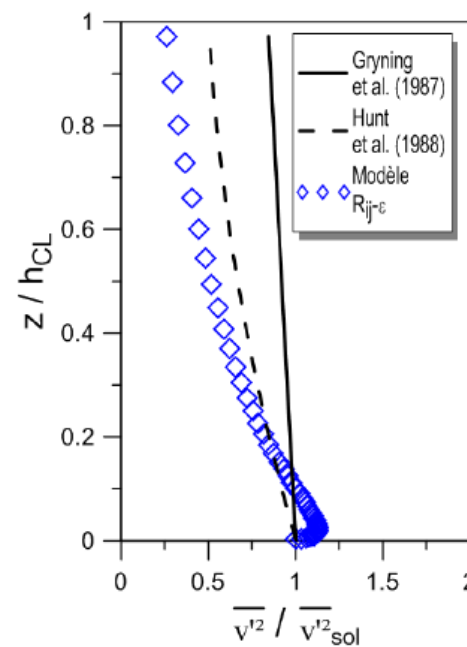
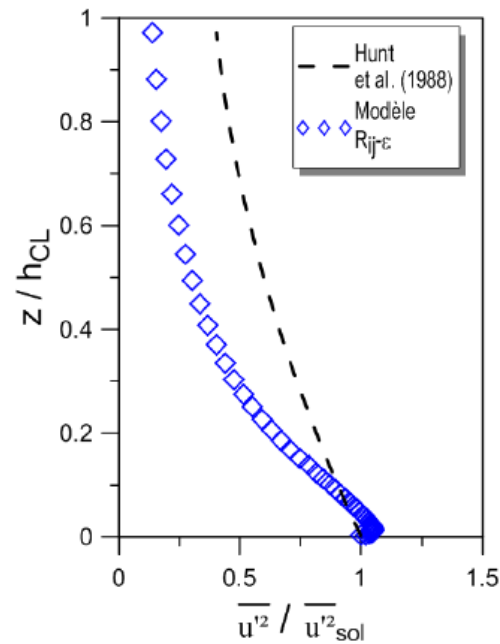


Stable conditions

# 4 – Validation of the model

## Results in stratified conditions

- **Turbulence profiles (normalized by the ground value):**
  - For example for  $U_g = 5\text{m}\cdot\text{s}^{-1}$ ,  $z_0 = 0.01\text{m}$ ,  $\varphi = 45^\circ$

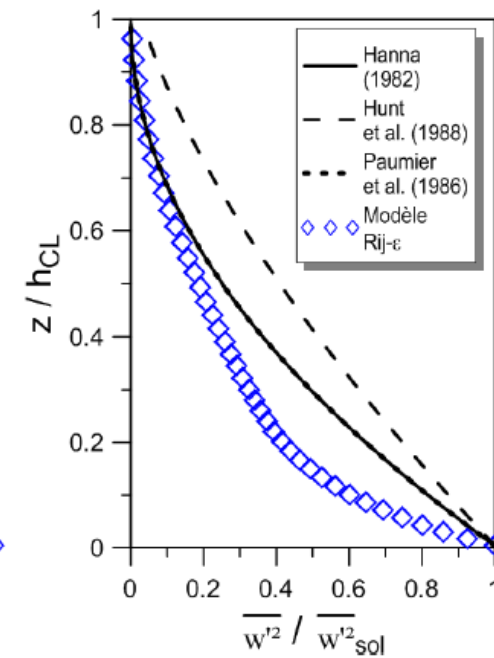
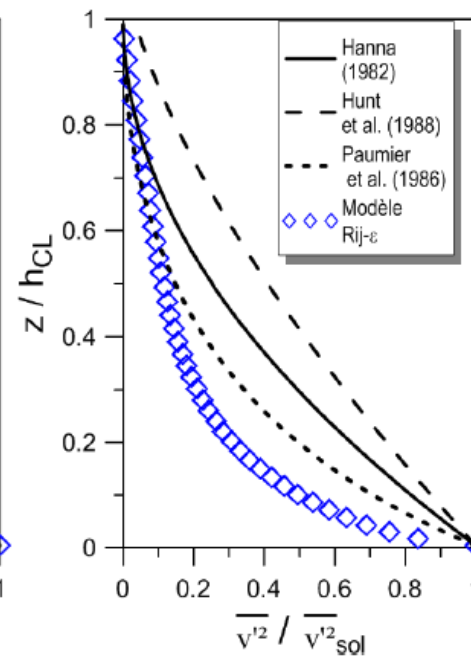
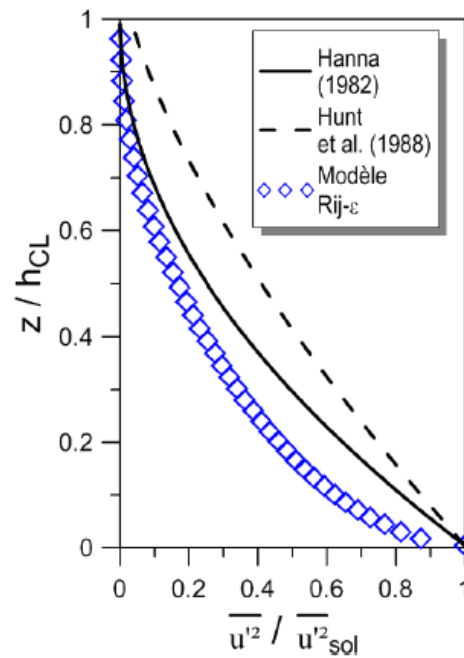


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Stable conditions

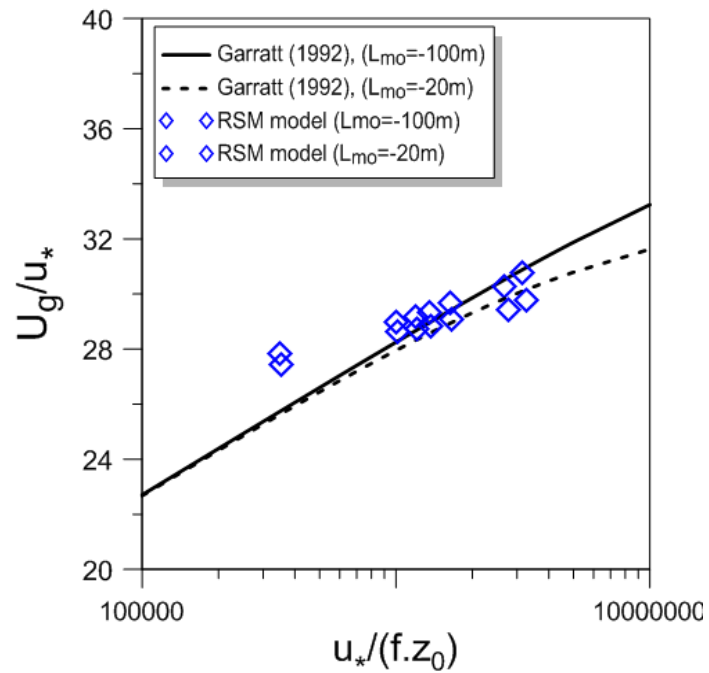
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## Results in stratified conditions

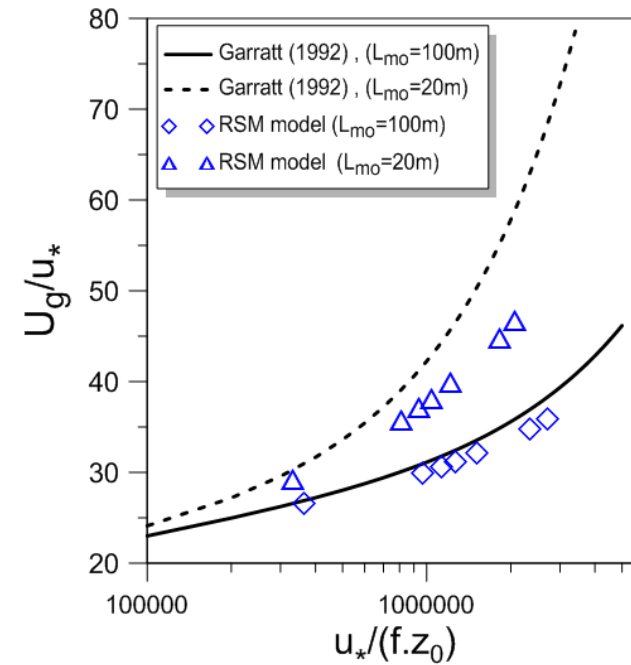
- Relation between  $U_g$  and  $u_*$ :

$$\frac{U_g}{u_*} = \frac{1}{\kappa} \sqrt{\left( \ln\left( \frac{u_*}{fz_0} \right) - B \right)^2 - A^2}$$

*A and B are dependent on stability, measured empirically*



Unstable conditions



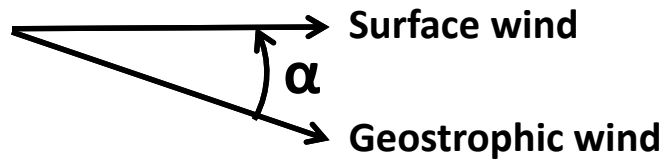
Stable conditions



# 4 – Validation of the model

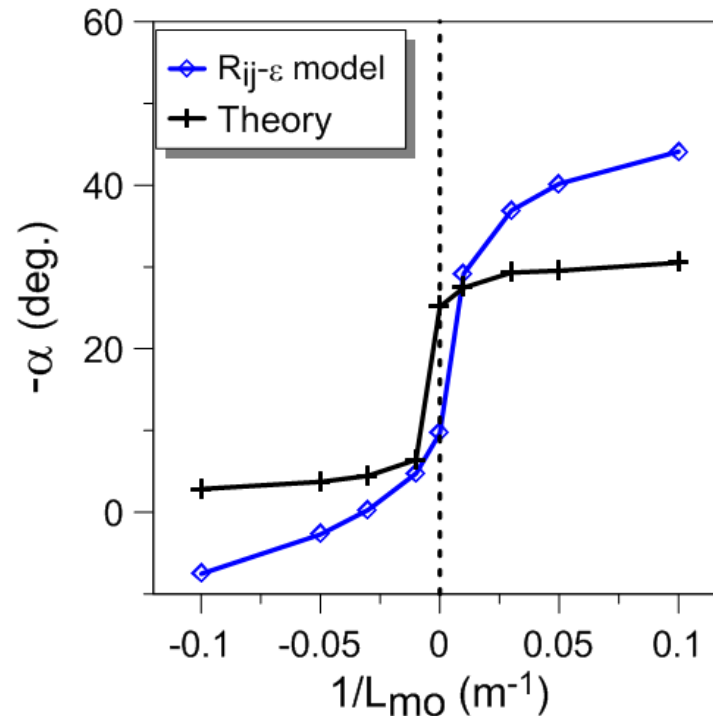
## Results in stratified conditions

- The Rossby similarity theory gives:



$$\sin(\alpha) = -\frac{A u_*}{\kappa U_g} \text{sign}(f)$$

*A is dependent on stability, measured empirically*







## 5 – Conclusions and perspectives

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- **Conclusions**
  - Development of a **new set of atmospheric constants** for the Reynolds Stress turbulence model
  - **1D simulations** of the Atmospheric Boundary Layer with this Reynolds Stress Model, **including Coriolis effects**
  - **Validation against empirical results**
- **Perspectives**
  - Validation in more complex configurations
  - Evaluation of the effect of anisotropy on dispersion modeling
  - Unified approach between “atmospheric” and “standard” RSM constants (see Poster H15-166)



**Thank you for your attention 😊**

**Questions ?**